

Basic concepts

1. Atmospheric **continuum**
2. Physical dimensions, variables and units
3. Conservation laws (momentum, mass, energy)
4. Eulerian and Lagrangian viewpoints



Source: Videezy.com

Concept of a “continuum”

The atmosphere (or fluid) is composed of molecules, but!

“The atmosphere can be regarded as **a continuous fluid medium, or continuum**. A point in the continuum is regarded as a volume element that is very small compared with the volume of atmosphere under consideration, but still contains a large number of molecules...

air parcel and air particle are both commonly used to refer to such a point” (*Intro. dyn. Met.*, Holton)

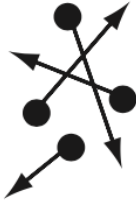
But, I'm not the first one who mentioned this concept. Augustin-Louis Cauchy (1798-1857) is the one.

Professor James Reed Holton (1938-2004)



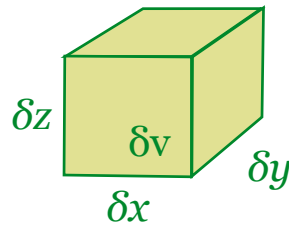
Concept of a “continuum”

Molecules



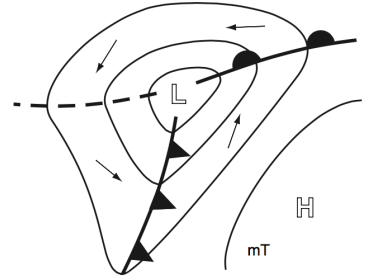
$l \sim 68 \text{ nm}$
(mean free path
under STP)

Control volume



$\delta x, \delta y, \delta z$

Phenomena of interest



$L \sim 5000 \text{ km}$

Figures from "Practical Meteorology (R. Stull)"

Concept of a “continuum”

“a volume element (control volume, air parcel) contains a **large number** of molecules”

$\gg \sim 2.8 \times 10^{19} \text{ molecules/cm}^3$

under standard temperature (273.15 K) and pressure (1013 hPa)

Physical quantities (**pressure, density, temperature, wind**) are **uniquely defined** at each point (air parcel) under the concept of atmospheric continuum.

International system of unit (SI unit)

- SI Base Units

Property	Unit	Name
Length	m	Meter
Mass	kg	Kilogram
Time	s	Second
Temperature	K	Kelvin

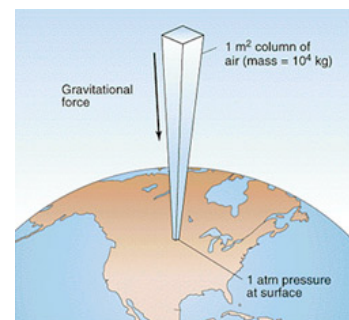
- Derived Units

Property	Unit	Name
Force	N (kg m s^{-2})	Newton
Pressure	Pa (N m^{-2})	Pascal
Energy	J (N m)	Joule
Power	W (J s^{-1})	Watt
Frequency	Hz (s^{-1})	Hertz

Atmos. variables : Pressure (P)

Easy definition (hydrostatic pressure)

Weight of air column
above the measurement point



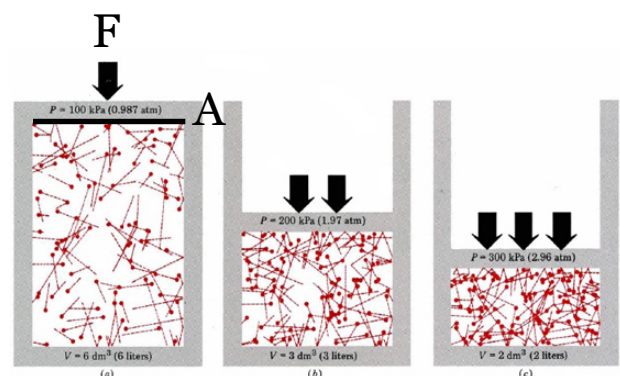
(from faculty.sdmiramar.edu/)

General definition

Force acting perpendicular to a unit surface ($P = F/A$)
(by randomly moving molecules that bounce off the surface
- Atmospheric pressure)

* Units: N/m^2 (Pascal, Pa)

* Standard sea-level pressure:
101325 Pa (1013.25 hPa)

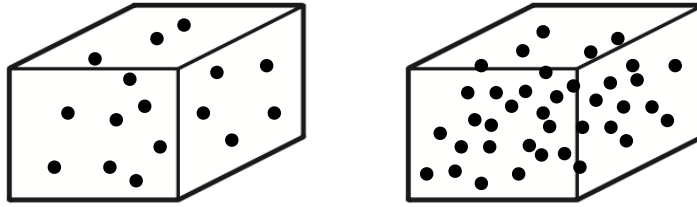


(from ChemEd DL)

Atmos. variables : Density (ρ)

Density is defined as mass (m) per unit volume (V):

$$\rho = m/V$$



- * Units: kg/m^3
- * Sea-level density at 288 K (15°C)
 $\rho_o = 1.225 \text{ kg/m}^3$
- * Gases such as air are compressible, air density can vary.
- * Atmospheric density decreases exponentially with height.

Atmos. variables : Temperature (T)

Simple concept

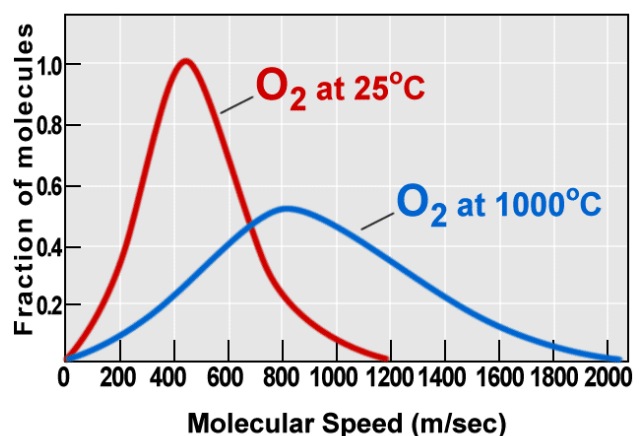
Measure of “hot” and “cold”

Kinematic temperature

Higher temperatures T are associated with greater average molecular speeds v

$$N \left[\frac{1}{2} m v^2 \right]_{\text{avg}} = \frac{3}{2} n R^* T$$

N number of molecules
 m mass of molecule (kg)
 n amount of molecules (in moles)
 R^* gas constant ($=8.314 \text{ J K}^{-1}\text{mol}^{-1}$)
($N/n = N_A$, Avogadro's number)



(from www.dlt.ncssm.edu)

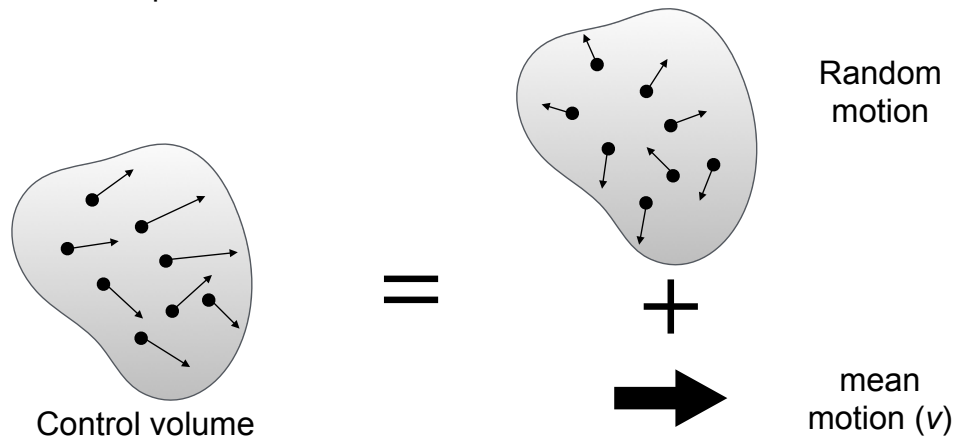
©NCSSM 2002

Atmos. variables : Wind (v)

Simple concept

“When a group of molecules move predominantly in the same direction, the motion is called **wind**” (Practical Meteorology)

Recall the concept of **continuum**



Note: difference from Kinematic temperature

When molecules move in random directions, the motion is associated with temperature (Practical Meteorology).

Ideal gas

“An **ideal gas** is a theoretical gas composed of many randomly moving point particles that do not interact except when they collide elastically... At normal conditions, most real gases (N₂, O₂, H₂, etc) behave like an ideal gas” (wikipedia)

Ideal gas law

$$PV = nR^*T$$

- P pressure (Pa)
- V volume (m³)
- T temperature (K)
- n amount of molecules (in moles)
- R^* gas constant (=8.314 J·K⁻¹mol⁻¹)

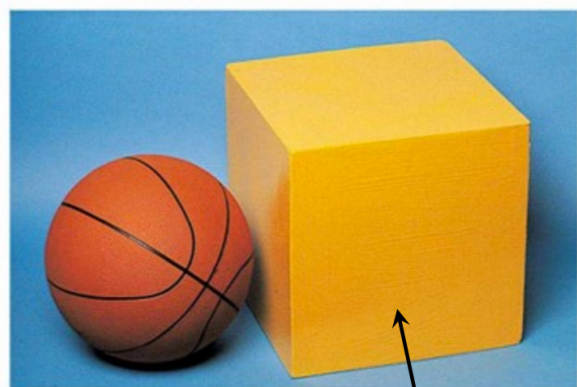


Photo by James Scherer. ©Houghton Mifflin Company. All rights reserved.

22.4 L
(22.4 x 10⁻³ m³)

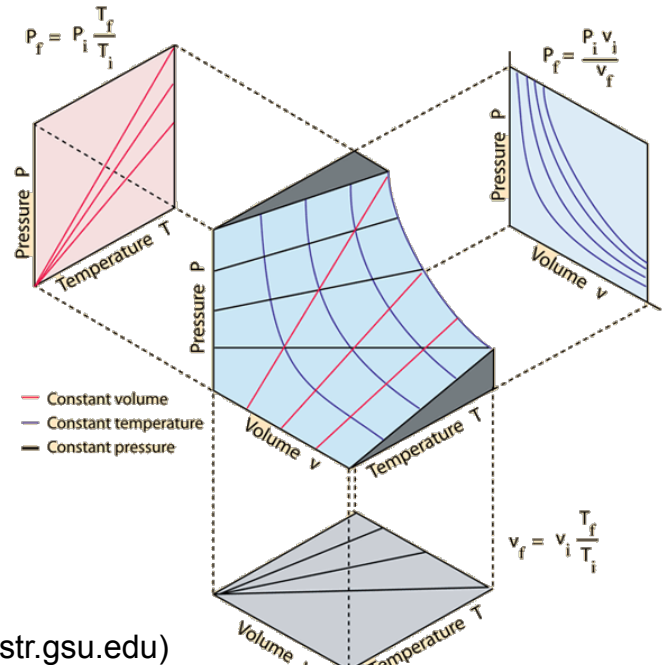
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(from hyperphysics.phy-astr.gsu.edu)

Ideal gas

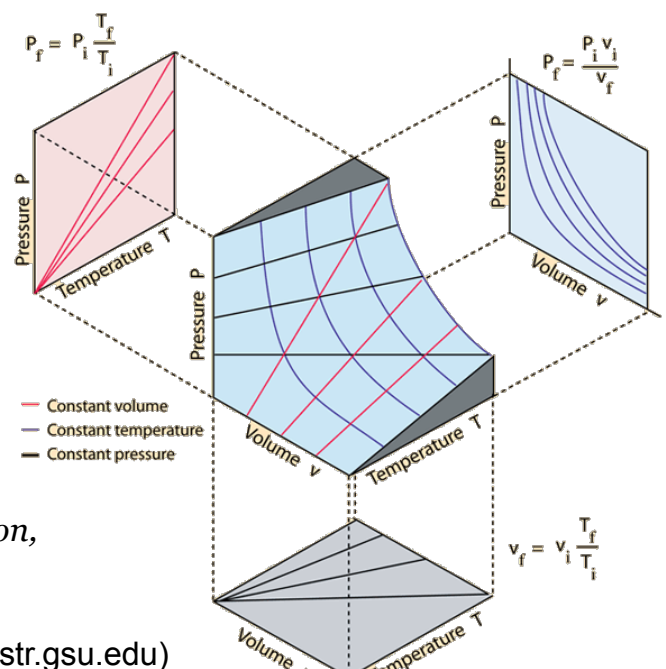
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Ideal gas law

$$PV = nR^*T$$

n amount of molecules (in moles)
 R^* gas constant ($=8.314 \text{ J}\cdot\text{K}^{-1}\text{mol}^{-1}$)

* Under standard temperature (273K) and pressure (1 atm, 1013.25 hPa) condition, 1 mole of ideal gas has volume of 22.4L



(from hyperphysics.phy-astr.gsu.edu)

Ideal gas law (equation of state)

Atmospheric version

$$P = \rho RT$$

P pressure (Pa)

ρ density (kg m^{-3})

T temperature (K)

R specific gas constant ($=287 \text{ J K}^{-1} \text{ kg}^{-1}$) of dry air - atmospheric version

Derivation

$$PV = nR^*T$$

n amount of molecules (in moles)

R^* gas constant ($=8.314 \text{ J K}^{-1} \text{ mol}^{-1}$)

M molar mass of dry air (kg mol^{-1})

$$P = nR^*T/V = nR^*(M/M)T/V = \underbrace{(nM/V)}_{\rho} \times \underbrace{(R^*/M)}_R \times T$$

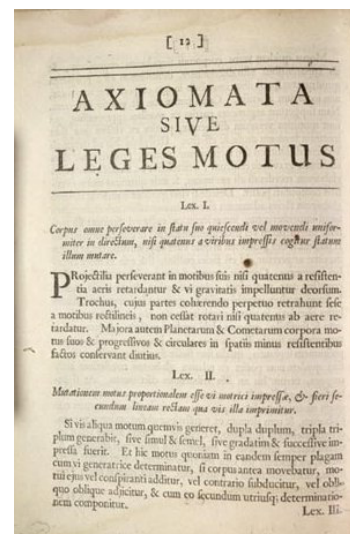
$$P = \rho RT$$

(where $R = R^*/M$)

Conservation of momentum

Newton's law of motion

1. When viewed in an inertial reference frame, an object either **remains** at rest or continues to move **at a constant velocity**, unless acted upon by a force.
2. The vector sum of the forces **F** on an object is equal to the mass m of that object multiplied by the acceleration vector **a** of the object:
F = ma.
3. When one body exerts a force on a second body, the second body simultaneously exerts a force equal in magnitude and opposite in direction on the first body.



Principia Mathematica, 1687
(obtained from wikipedia)

Conservation of momentum

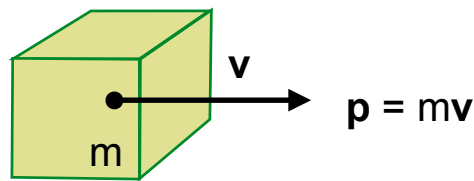
Momentum (**p**)

The quantity of motion of a moving body, measured as a product of its mass and velocity.

$$\mathbf{p} = m\mathbf{v} \quad (\text{bold-face means a vector quantity})$$

* Units: kg m/s (N s)

$$* \mathbf{F} = \Delta\mathbf{p}/\Delta t \quad [\mathbf{F} = \Delta(m\mathbf{v})/\Delta t \rightarrow \mathbf{F} = m\mathbf{a}]$$



Conservation of momentum

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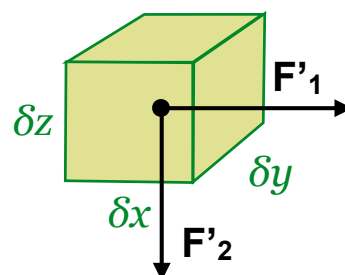
$$* \int_0^t \vec{\mathbf{F}} dt = \int_0^t m\vec{\mathbf{a}} dt \rightarrow \int_0^t \vec{\mathbf{F}} dt = m\vec{\mathbf{v}} \Big|_{t=0}^{t=t'} = \vec{\mathbf{p}} \Big|_{t=0}^{t=t'}$$

Atmospheric version

$$\underbrace{(\rho\delta x\delta y\delta z)}_m \underbrace{\frac{D\vec{\mathbf{v}}}{Dt}}_a = \underbrace{\vec{\mathbf{F}}_1 + \vec{\mathbf{F}}_2 + \dots}_{\Sigma \mathbf{F}}$$

$$\frac{D\vec{\mathbf{v}}}{Dt} = \frac{1}{\rho} \frac{\vec{\mathbf{F}}_1 + \vec{\mathbf{F}}_2 + \dots}{\delta x\delta y\delta z} = \frac{1}{\rho} \vec{\mathbf{F}}_1' + \frac{1}{\rho} \vec{\mathbf{F}}_2' + \dots$$

Force per unit volume

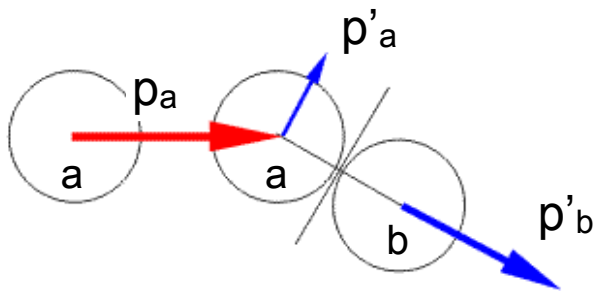


Conservation of momentum

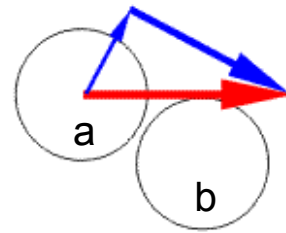
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(from wikipedia, [momentum movie](#))



$$\mathbf{p}_a + \mathbf{p}_b = \mathbf{p}'_a + \mathbf{p}'_b$$

Conservation of mass

Eulerian view (fixed point)

Rate of mass change in a fixed volume must equal the rate of net mass influx through the walls

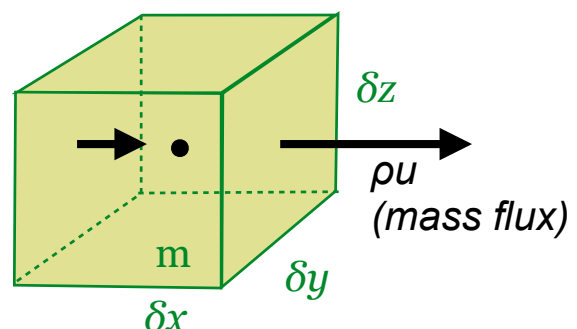
$$\begin{aligned} \frac{\partial(m)}{\partial t} &= \left[-\frac{\partial(\rho u \delta y \delta z)}{\partial x} \left(\frac{\delta x}{2} \right) \right] + \left[\frac{\partial(\rho u \delta y \delta z)}{\partial x} \left(-\frac{\delta x}{2} \right) \right] + \dots \\ &= -\frac{\partial(\rho u)}{\partial x} \delta x \delta y \delta z - \frac{\partial(\rho v)}{\partial y} \delta x \delta y \delta z - \frac{\partial(\rho w)}{\partial z} \delta x \delta y \delta z \end{aligned}$$

$$\begin{aligned} \frac{\partial \rho}{\partial t} &= -\frac{\partial(\rho u)}{\partial x} - \frac{\partial(\rho v)}{\partial y} - \frac{\partial(\rho w)}{\partial z} \\ \frac{\partial \rho}{\partial t} + \frac{\partial(\rho u)}{\partial x} + \frac{\partial(\rho v)}{\partial y} + \frac{\partial(\rho w)}{\partial z} &= 0 \end{aligned}$$

$$\frac{\partial \rho}{\partial t} + \nabla \cdot (\rho \vec{v}) = 0$$

popular form

$$m = \rho \delta x \delta y \delta z$$



Conservation of energy

First law of thermodynamics

For a closed system,

$$\Delta U = Q - W \quad (\text{or } Q = \Delta U + W)$$

*Heat supplied
from outside*

*change in
Internal energy*

*Work done
by the system*

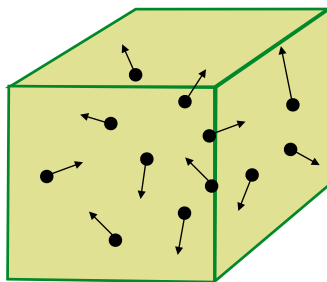
Atmospheric version (ideal gas, per unit mass)

$$c_v \Delta T + p \Delta \alpha = q$$

$$c_v \frac{DT}{Dt} + p \frac{D\alpha}{Dt} = J$$

$$c_p \frac{DT}{Dt} - \alpha \frac{Dp}{Dt} = J$$

popular form



$$\begin{aligned} \Delta U/m &= c_v T \\ &= (5/2) RT \end{aligned}$$

$$c_v = 717.6 \text{ J K}^{-1} \text{ kg}^{-1}$$

(specific heat at constant volume)

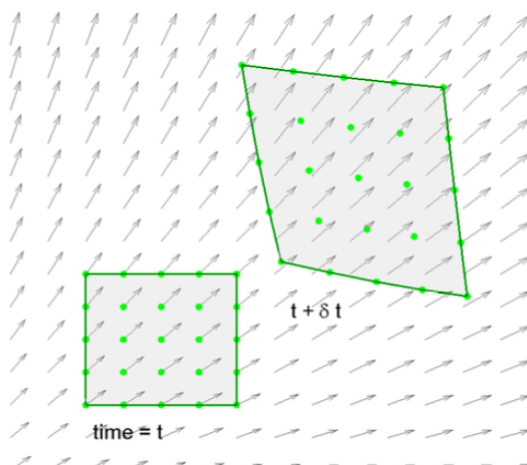
Eulerian vs. Lagrangian

Eulerian description

Describe properties (p, T, ρ, \mathbf{v}) of fluid flow **at a fixed point** or volume (easy for math description and measurements)

Lagrangian description

Describe properties (p, T, ρ, \mathbf{v}) of fluid flow **following a parcel** or volume (easy to think; ' $\mathbf{F} = m\mathbf{a}$ ' works with this concept)



NSF Fluid Mech lecture

(concept ~1.5min, equation ~12min)

[https://www.youtube.com/watch?](https://www.youtube.com/watch?v=mdN80Okx2ko)

[v=mdN80Okx2ko](https://www.youtube.com/watch?v=mdN80Okx2ko)

J. Price 2006
(Woods Hole Oceanographic Institution)

Material derivative (total derivative)

Material derivative, $D(\)/Dt$

describes the time rate of change of some physical quantity (p, T, ρ, \mathbf{v}) for a material element (**Lagrangian**) subjected to a space-and-time-change.

$$\frac{DT}{Dt} \equiv \frac{\partial T}{\partial t} + u \frac{\partial T}{\partial x} + v \frac{\partial T}{\partial y} + w \frac{\partial T}{\partial z}$$

Lagrangian $\equiv \frac{\partial T}{\partial t} + \vec{v} \cdot \nabla T$ *Eulerian*

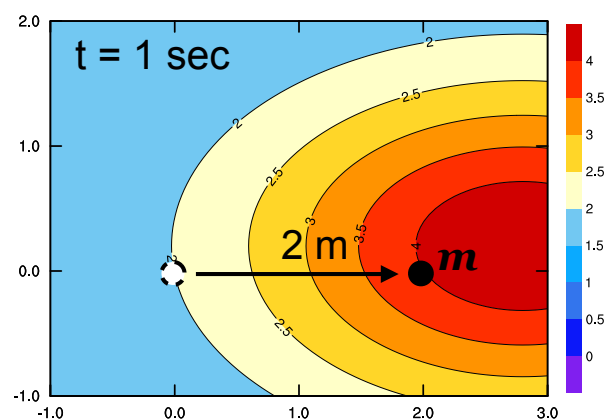
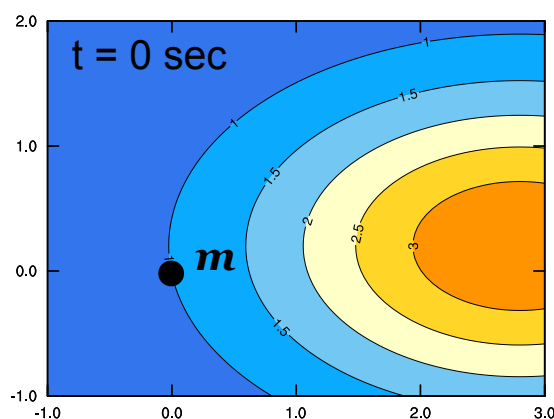
$\partial(\)/\partial t$ partial derivative
(measures local change, **Eulerian**)

Material derivative (total derivative)

Material derivative, $D(\)/Dt$

$$\frac{DT}{Dt} \equiv \frac{\partial T}{\partial t} + u \frac{\partial T}{\partial x} + v \frac{\partial T}{\partial y} + w \frac{\partial T}{\partial z}$$

3 K/s 1 K/s $2 \text{ m/s} \times (2 \text{ K})/(2 \text{ m})$

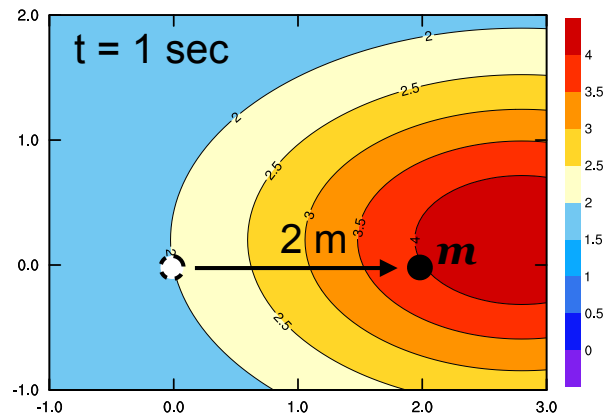
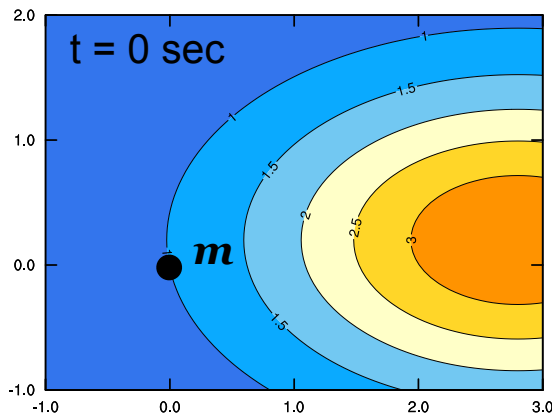


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Material derivative, $D(\)/Dt$

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$\uparrow \quad \quad \uparrow \quad \quad \uparrow$
 $3 \text{ K/s} = 1 \text{ K/s} + 2 \text{ K/s}$



Conservation of mass

Lagrangian view (following material)

Mass of a material volume should be conserved with motion.

$$\frac{D(m)}{Dt} = 0 \quad m = \rho \delta x \delta y \delta z$$

$$\frac{D(\rho \delta x \delta y \delta z)}{Dt} = \frac{D(\rho)}{Dt} \delta x \delta y \delta z + \rho \frac{D(\delta x \delta y \delta z)}{Dt} = 0$$

$$\frac{D(\rho)}{Dt} + \rho \frac{1}{\delta x \delta y \delta z} \frac{D(\delta x \delta y \delta z)}{Dt} = 0$$

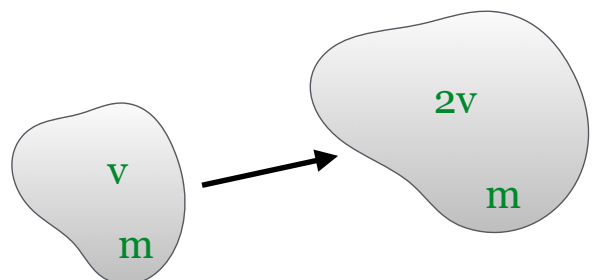
$$\left(\text{using } \lim_{\delta x \rightarrow 0} \frac{1}{\delta x} \frac{D\delta x}{Dt} = \frac{\partial u}{\partial x} \right)$$

$$\frac{D\rho}{Dt} + \rho \nabla \cdot \vec{v} = 0$$

popular form

$$\frac{\partial \rho}{\partial t} + \nabla \cdot (\rho \vec{v}) = 0$$

can change into Eulerian form



References

- Introduction to Dynamic Meteorology (J. R. Holton)
- Practical Meteorology (R. Stull, 2015)
available in web (https://www.eoas.ubc.ca/books/Practical_Meteorology/)