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# Impact of North Atlantic-East Asian teleconnections on extremely high January PM<sub>10</sub> cases in Korea<sup>☆</sup>

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## ABSTRACT

In this study, we investigated the daily variability of PM<sub>10</sub> concentrations in January in Korea during the past 19 years (2001–2019), as well as the associated atmospheric circulation patterns. The daily PM<sub>10</sub> concentrations were classified into three cases: low (L; < 50 µg/m<sup>3</sup>), high (H; 50–100 µg/m<sup>3</sup>), and extremely high (EH; ≥ 100 µg/m<sup>3</sup>). We found that the strength of the East Asian winter monsoon influenced the PM<sub>10</sub> variability in the L and H cases. However, the EH cases were strongly influenced by the rapid growth of barotropic warming (anticyclonic anomaly) over the eastern North Atlantic and Northern Europe (ENE), and the stationary Rossby waves grew rapidly over Eurasia within only four days. Analysis of the quasi-geostrophic geopotential tendency budget revealed that the anticyclonic anomaly over the ENE was enhanced by vorticity advection. Linear baroclinic model experiments confirmed that vorticity forcing over the ENE induces favorable atmospheric conditions for the occurrence of EH PM<sub>10</sub> events in East Asia. As a result, the PM<sub>10</sub> concentration sharply increased sharply by approximately three times over four days. This study suggests that understanding atmospheric teleconnections between the ENE and East Asia can effectively predict the occurrence of EH PM<sub>10</sub> events in Korea, helping to reduce the human health risks from atmospheric pollution.

## 1. Introduction

Over the last decade, high particulate matter (PM) concentration events in the winter season have occurred most frequently in East Asia (Wang et al., 2015; Wang and Chen, 2016; Cai et al., 2017; Kim et al., 2019). These events have increased the risk of cardiovascular and respiratory diseases (Brunekreef and Holgate, 2002; Nel, 2005) in China and Korea. A prime example was the high PM concentration event in January 2013 in China and Korea, referred to as *airpocalypse*, which caused significant damage (Park et al., 2013; Wang et al., 2014; Zou et al., 2017; Koo et al., 2018). In Beijing, this high PM episode in 2013 caused 476 acute deaths, equivalent to ~180 million dollars in damage or ~0.76% of the GDP (Du and Li, 2016). Since the early 2000s, the Korean government has continually reduced emissions from factories and various transportation types through the “National Environmental Comprehensive Plan” (Korea Ministry of Environment, 2015; Kim et al., 2016; Lee et al., 2020b; Bae et al., 2021), helping to decrease the

national atmospheric PM concentration (Lee et al., 2020b; Oh et al., 2015). Nevertheless, high PM concentration events are stronger and more frequent in January (Jung et al., 2019; Kwon et al., 2020).

PM in which 50% of particles have an aerodynamic diameter less than 10 µm (PM<sub>10</sub>) is mainly emitted by factories, cars, and power plants (Chen et al., 2013; Kim et al., 2016; Ryou et al., 2018; Rai et al., 2021). In particular, high PM<sub>10</sub> pollution in Korea and China is affected by combined factors such as emission sources (Kim et al., 2016; Kim et al., 2017a, b), inflow from neighboring countries (Seo et al., 2018; Hur et al., 2021), and meteorological conditions (Dotse et al., 2016; Guan et al., 2017). PM<sub>10</sub> emission sources can be classified into natural and anthropogenic (Kim et al., 2016; Seo et al., 2017), which are both affected by distinct sources and atmospheric circulation patterns (Lee et al., 2018; Jung et al., 2019). Yellow dust, a natural pollutant, is mainly transported by winds from the desert regions of northern China and Mongolia (Lee and Kim, 2018). However, PM<sub>10</sub> is primarily caused by local emissions in South Korea, transport from China and North Korea

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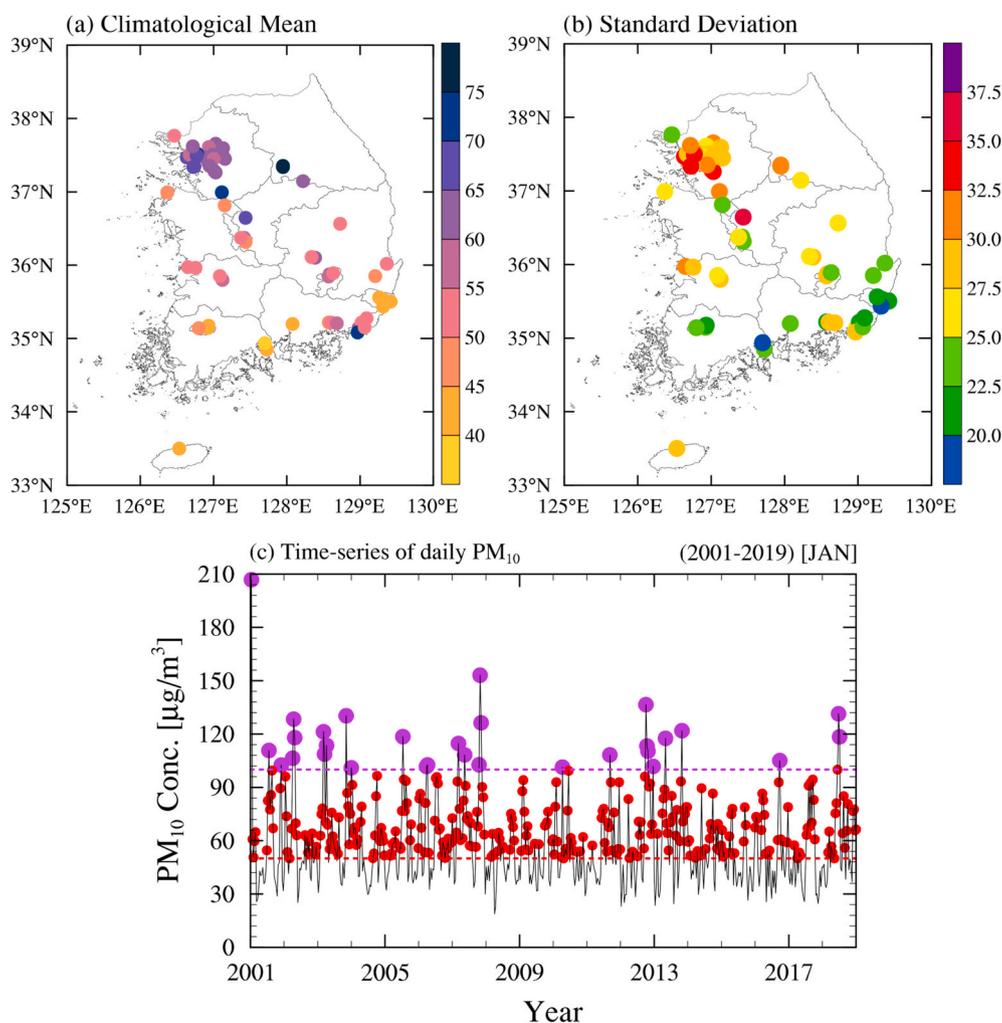
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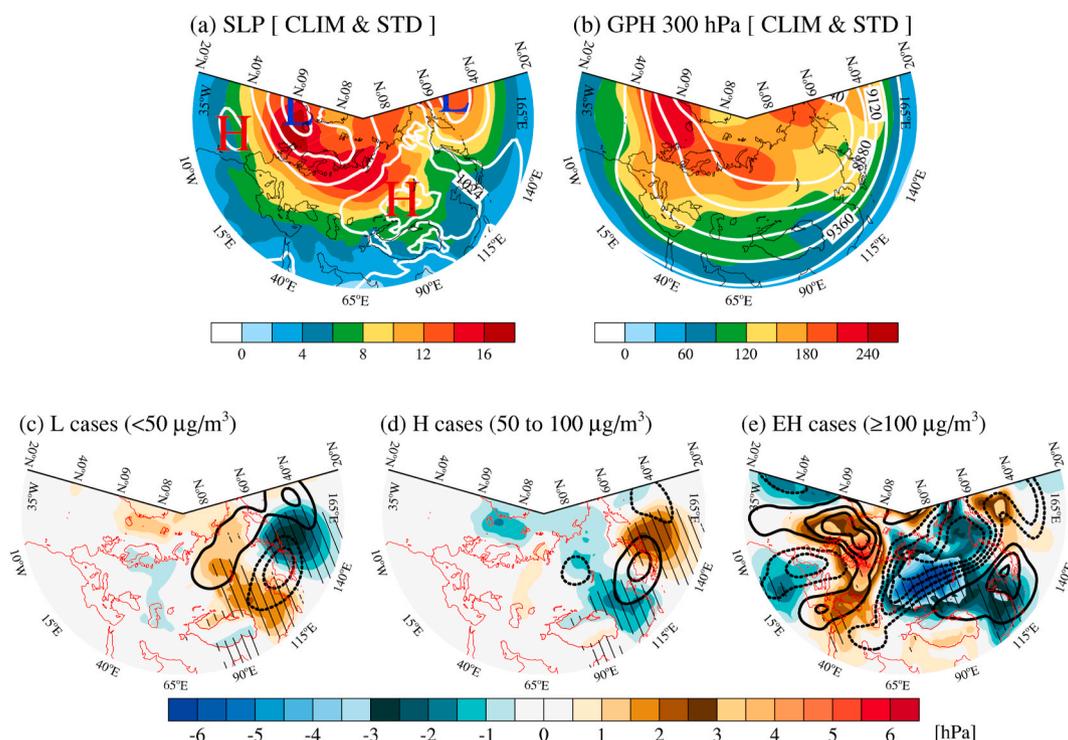
**Fig. 1.** The spatial distribution of (a) climatological mean, (b) standard deviation, and (c) the time-series of January  $PM_{10}$  concentrations over 19 years (2001–2019). The red and purple circles in (c) indicate H cases ( $50\text{--}100\ \mu\text{g}/\text{m}^3$ ) and EH cases ( $\geq 100\ \mu\text{g}/\text{m}^3$ ), respectively. The red and purple dashed lines indicate 50 and  $100\ \mu\text{g}/\text{m}^3$ , respectively. (For interpretation of the references to colour in this figure legend, the reader is referred to the Web version of this article.)

(Lee et al., 2013; Kim et al., 2016), and weakening dispersion by stagnant atmospheric circulation patterns (Seo et al., 2018; Shi et al., 2019; Oh et al., 2020). The stagnant atmospheric circulation pattern that influences the high  $PM_{10}$  concentration events in East Asia reduces the dispersion of  $PM_{10}$  pollutants by strengthening the horizontal and vertical atmospheric stability (Lee et al., 2011; Zou et al., 2017; Kim et al., 2019; Kwon et al., 2020).

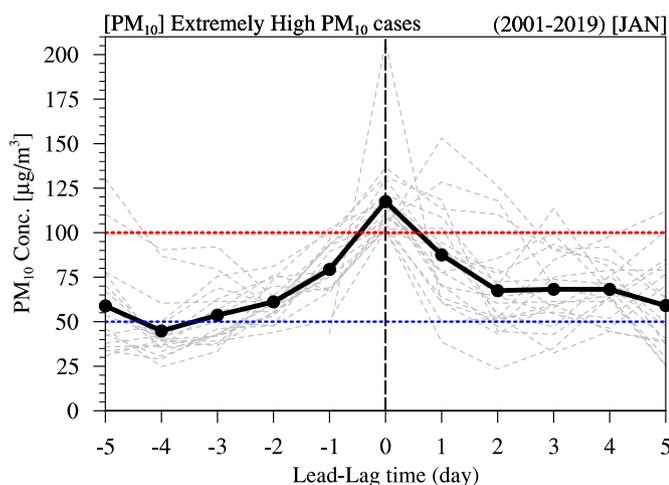
Stagnant atmospheric circulation patterns are affected by various teleconnections. Previous studies have suggested that a weakening of the East Asian winter monsoon (EAWM) induces stagnant atmospheric circulation, a much shallower boundary layer height, and stable atmospheric stratification, resulting in increased aerosols and  $PM_{10}$  concentrations in China and Korea (Niu et al., 2010; Kim et al., 2016; Jeong and Park, 2017). In addition, the Siberian high pressure and Aleutian low pressure are the dominant surface pressure systems in East Asia that determine the wind direction and wind speed during the winter season associated with the EAWM; the variability in these systems can therefore lead to high  $PM_{10}$  pollution in East Asia (Oh et al., 2018; Kim et al., 2019). A decrease in Arctic sea ice concentration and Eurasian snow cover has been suggested to weaken the Siberian high pressure, which suppresses the northerly winds and induces a stagnant atmospheric circulation pattern in China and the Korean Peninsula, causing high  $PM_{10}$  concentration events during the winter season (Wang et al., 2015;

Zou et al., 2017; Kim et al., 2019). Lee et al. (2020b) found that the weakening of the Ural blocking since 2014 suppressed cold air flows from the north, causing favorable atmospheric conditions for poor air quality in East Asia. In addition, it was reported that the El Niño–Southern Oscillation (ENSO) alters the haze days by modulating winter precipitation in northern China (He et al., 2019). The El Niño event reduces wet deposition and induces haze events. The  $PM_{10}$  concentration on the Korean Peninsula is also closely related to the precipitation change caused by ENSO (Wie and Moon, 2017). However, the relationship between North Atlantic forcing and the high  $PM_{10}$  pollution in East Asia is relatively unknown.

We hypothesize that remote forcing over the North Atlantic may play an important role in triggering extremely high  $PM_{10}$  events in East Asia. Knowing this information in advance can minimize the health impacts caused by extremely high  $PM_{10}$  episodes. Therefore, this study examined the teleconnection between North Atlantic forcing and stagnant atmospheric patterns associated with the occurrence of high  $PM_{10}$  events during January 2001–2019 in Korea. Atmospheric variability is typically higher in January, mainly due to the influence of a strong or weak EAWM (Zhao et al., 2018). However, as the impact of yellow dust on  $PM_{10}$  concentrations is relatively small and high  $PM_{10}$  events are most frequent (Jung et al., 2019; Kwon et al., 2020; Ku et al., 2021), January is considered an ideal month for studying the impact of atmospheric



**Fig. 2.** (a–b) Climatological mean (white contour) and standard deviation (shading) of (a) sea level pressure (SLP; contour interval of 4 hPa) and (b) upper-level (300 hPa) geopotential height (GPH; contour interval of 240 m). (c–e) Composite anomalies of the January SLP (shading) and GPH at 300 hPa (contour interval of 20 m; dashed contours are negative) for L (<50 μg/m<sup>3</sup>), H (50–100 μg/m<sup>3</sup>), and EH (≥100 μg/m<sup>3</sup>) cases, respectively. The hatched areas indicate a 10% significance level for SLP.



**Fig. 3.** Lead-lag time series of PM<sub>10</sub> concentrations from five days before to five days after the occurrence of EH cases. The black solid line indicates the mean PM<sub>10</sub> concentration, and the gray dashed lines indicate the PM<sub>10</sub> concentration of each EH case. The red and blue dashed lines indicate PM<sub>10</sub> concentrations of 50 μg/m<sup>3</sup> and 100 μg/m<sup>3</sup>, respectively. (For interpretation of the references to colour in this figure legend, the reader is referred to the Web version of this article.)

circulation on PM<sub>10</sub> variability. We employed statistical methods to define extremely high PM<sub>10</sub> cases and identify atmospheric circulation patterns associated with extreme cases. In addition, we applied a dynamical method and modeling to understand the teleconnection mechanisms between the anticyclonic anomaly over the North Atlantic and the stagnant atmospheric conditions in Korea associated with extremely high PM<sub>10</sub> episodes.

## 2. Methods

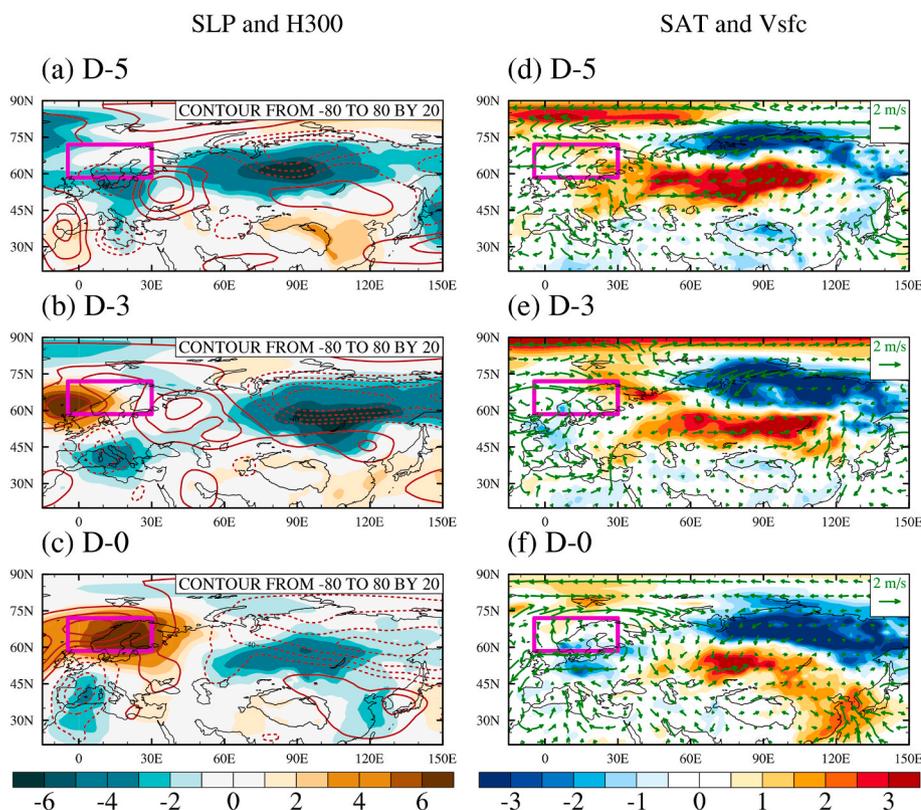
### 2.1. Observation and reanalysis data

In this study, we used daily PM<sub>10</sub> data collected over 19 years (2001–2019) from 68 observational stations provided by the National Institute of Environmental Research (NIER), Korea (<https://www.airkorea.or.kr>) (Fig. 1a). To ensure the statistical reliability of the data, 68 stations, which possess more than 70% of the January data during the analysis period, were used in this study. According to Yellow dust day information provided by Korean Meteorological Administration (KMA), we excluded Yellow dust days from the analysis. Thus, PM<sub>10</sub> data do not represent Yellow dust. Daily data were obtained by averaging the hourly PM<sub>10</sub> concentration data.

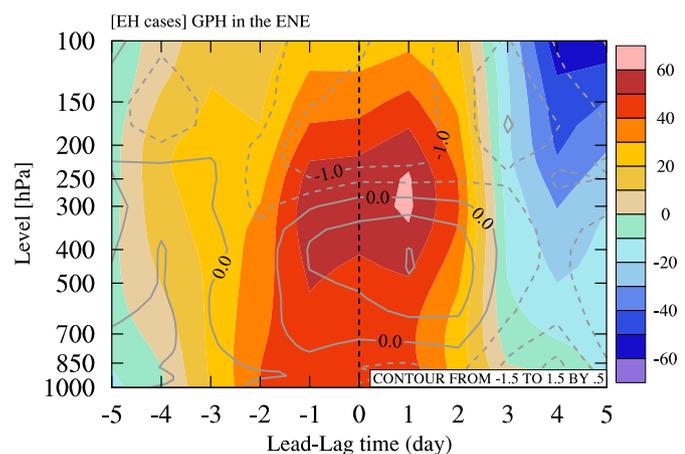
ERA-interim reanalysis data provided by the European Centre for Medium-Range Weather Forecasts (ECMWF) were used to investigate atmospheric circulation patterns associated with high PM<sub>10</sub> concentrations in Korea. The horizontal resolution of the data was 1.5° × 1.5° with 37 pressure levels. Our analysis of atmospheric circulations focused on the lower- (1000 and 850 hPa) and upper-level (300 hPa) geopotential height (GPH). The main variables used in the analysis were surface air temperature (SAT), sea level pressure (SLP), surface pressure (SP), GPH, air temperature (TMP), relative humidity (RH), specific humidity (SH), relative vorticity (RV), planetary boundary layer height (PBLH), horizontal wind (U- and V-wind), surface horizontal wind (Vsfc), and vertical velocity (W-wind).

### 2.2. Classification and composite analysis

In this study, we used the median and the 95th percentile values to classify the January PM<sub>10</sub> concentrations in Korea during the past 19 years. The daily PM<sub>10</sub> concentration were divided into three classes. In addition, we applied a composite analysis to determine the temporal evolution of daily weather patterns associated with high PM<sub>10</sub> pollution



**Fig. 4.** Composite anomalies of January (a–c) sea level pressure (SLP; units: hPa; shading) and upper-level (300 hPa) geopotential height (GPH; units: m; contour interval of 20 m, dashed lines are negative), and (d–f) surface air temperature (SAT; units: K; shading) and surface winds (Vsf; vectors) before and during the occurrence of EH cases. The pink boxes indicate the eastern North Atlantic and Northern Europe (ENE; 58.5 °N to 72 °N, 4.5 °W to 30 °E). (For interpretation of the references to colour in this figure legend, the reader is referred to the Web version of this article.)



**Fig. 5.** Vertical distribution of the composite anomaly of geopotential height (GPH; shading; units: m) and temperature (TMP; contour interval of 0.5 K; units: K; dashed contours are negative) in the eastern North Atlantic and Northern Europe (ENE; 58.5 °N to 72 °N, 4.5 °W to 30 °E) before and after the occurrence of EH cases. The black dashed line indicates the start day of EH cases.

in Korea.

### 2.3. Wave activity flux analysis

The three-dimensional wave activity flux (WAF) was calculated according to the formula of Takaya and Nakamura (2001). The WAF is useful diagnostics for Rossby wave’s energy propagation and related teleconnection. Rossby waves in the upper troposphere are associated with large-scale meanders of the jet stream, that are important in modulating atmospheric circulation patterns in the troposphere,

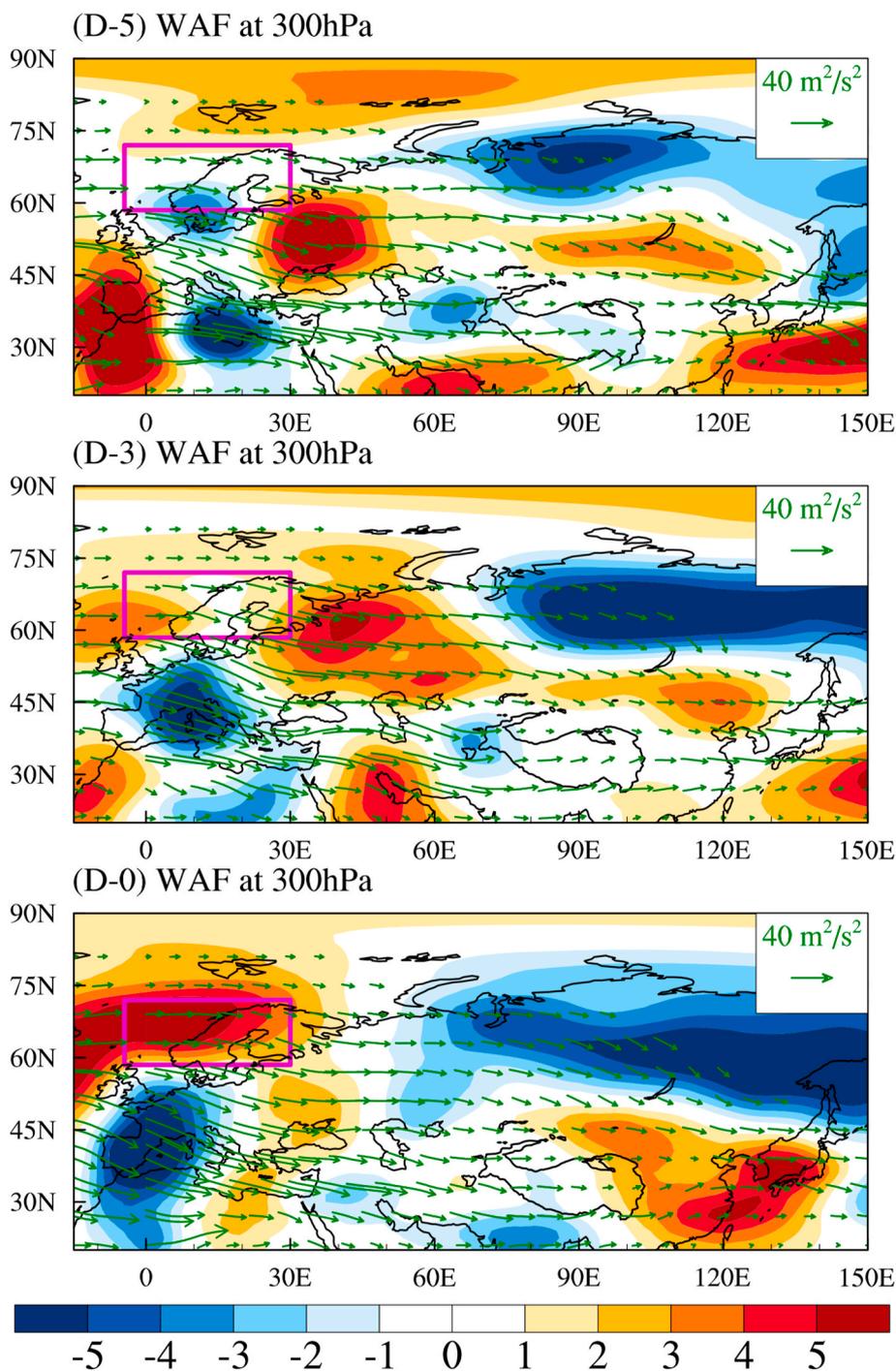
including the planetary boundary layer. Therefore, we analyzed the WAF to confirm the stationary Rossby waves from the North Atlantic forcing, which is related to the high PM<sub>10</sub> cases in Korea, as described in section 3.2.

### 2.4. Quasi-geostrophic geopotential tendency budget

The quasi-geostrophic geopotential tendency (QG tendency) equation was calculated following the method by Holton (2004). The QG tendency budget analysis is a helpful tool for quantitatively evaluating the contribution of dynamical processes (three forcing terms) to the temporal evolution of stationary high pressure (Hwang et al., 2020). The three forcing terms of the QG tendency equation are vorticity advection ( $F_{vort}$ ), temperature advection ( $F_{heat}$ ), and diabatic heating ( $F_{diab}$ ). The total forcing ( $F_{ALL}$ ) is defined as the sum of  $F_{vort}$ ,  $F_{heat}$ , and  $F_{diab}$ . In this study, the cause of the rapidly developed anticyclonic anomaly in the North Atlantic that affects the high PM<sub>10</sub> pollution in Korea (as described later) was analyzed through each forcing of the QG tendency.

### 2.5. Linear baroclinic model description

A linear baroclinic model (LBM) developed by Watanabe and Kimoto (2000) was used to confirm the effect of forcing on the change in GPH and to examine the steady atmospheric response to forcing (e.g., vorticity and heating) over targeted regions. The LBM used in this study had a dry dynamical core and is based on primitive equations linearized about a basic state on a sphere. The LBM consisted of a horizontal resolution of the T42 Gaussian grid and 20 vertical levels of sigma coordinates. The T42 is a 128 × 64 regular longitude/latitude global horizontal grid (approximately 2.8° resolution). The primary forces in the LBM are diabatic heating, vorticity, and divergence. In this study, we used the diagnosed forcing fields derived from reanalysis data. Further details of the model equations can be found in Watanabe and Kimoto (2000). The model experiment was integrated for 50 days, and the results are presented as the average of 10 days after reaching a steady



**Fig. 6.** Composite anomalies of the January quasi-geostrophic stream function (shading,  $10^6 \text{ m}^2/\text{s}$ ) and wave activity flux (WAF; vector,  $\text{m}^2/\text{s}^2$ ) at 300 hPa before and during the occurrence of EH cases. The pink boxes indicate the eastern North Atlantic and Northern Europe (ENE;  $58.5^\circ \text{N}$  to  $72^\circ \text{N}$ ,  $4.5^\circ \text{W}$  to  $30^\circ \text{E}$ ). (For interpretation of the references to colour in this figure legend, the reader is referred to the Web version of this article.)

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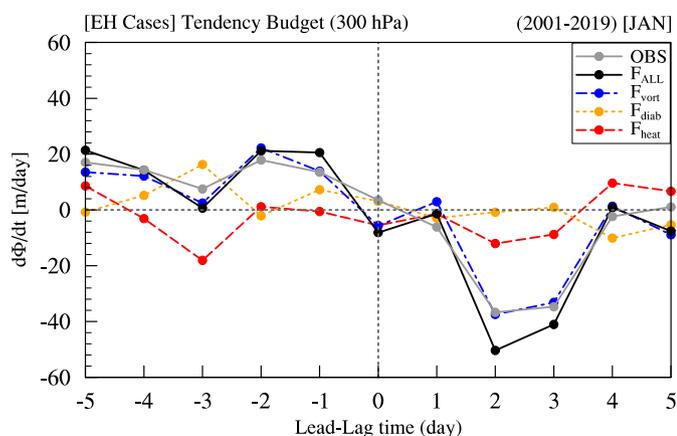
### 3. Results

#### 3.1. Characteristics of $\text{PM}_{10}$ concentration and related atmospheric circulation patterns

The mean and standard deviation of  $\text{PM}_{10}$  concentrations in South Korea were approximately  $56.9$  and  $22.5 \mu\text{g}/\text{m}^3$  in January, respectively (Fig. 1). The  $\text{PM}_{10}$  concentration range during this period was approximately  $18.8$ – $206.8 \mu\text{g}/\text{m}^3$ . In particular, the mean and standard

deviation of  $\text{PM}_{10}$  concentrations were high in Seoul and Gyeonggi-do, the metropolitan areas of Korea. Over the past 19 years, the median daily  $\text{PM}_{10}$  concentration in January was approximately  $52.0 \mu\text{g}/\text{m}^3$ , and the 95th percentile value was approximately  $100.9 \mu\text{g}/\text{m}^3$ . Based on this information, we divided the daily  $\text{PM}_{10}$  concentrations into three classes: low (L;  $< 50 \mu\text{g}/\text{m}^3$ ), high (H;  $50$ – $100 \mu\text{g}/\text{m}^3$ ), and extremely high (EH;  $\geq 100 \mu\text{g}/\text{m}^3$ ) (Fig. 1c). The number of H (EH) cases was 287 (30) over the past 19 years, corresponding to 48.7% (5%) of the total number of days (Fig. 1c).

To understand the atmospheric circulation patterns according to the classified  $\text{PM}_{10}$  concentrations, we investigated the composite patterns



**Fig. 7.** Time series of geopotential height tendency in the eastern North Atlantic and Northern Europe (ENE; 58.5 °N to 72 °N, 4.5 °W to 30 °E) at 300 hPa. The gray line indicates the tendency obtained from reanalysis data. The black, red, orange, and blue lines indicate the quasi-geostrophic geopotential tendencies induced by the total forcing, temperature advection, diabatic heating, and vorticity forcing, respectively. (For interpretation of the references to colour in this figure legend, the reader is referred to the Web version of this article.)

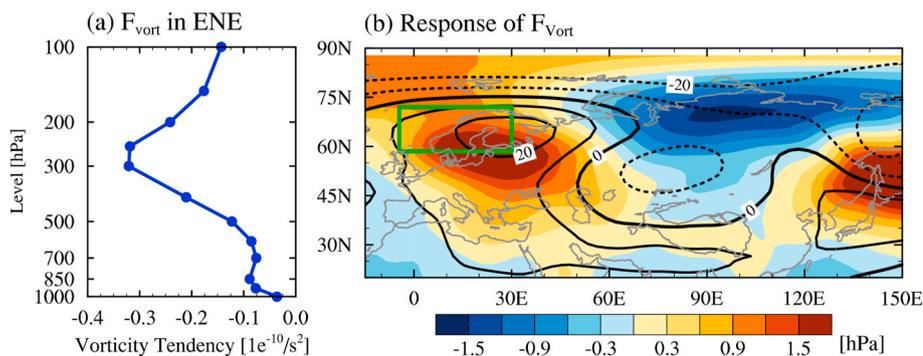
of each case (Fig. 2). The climatological mean distribution of SLP in January showed that the Siberian high pressure and Aleutian low pressure systems dominated in East Asia, and the Icelandic low pressure and Azores high pressure were predominant in the North Atlantic (Fig. 2a). The SLP variability was high in the North Atlantic Ocean, Kara-Barents seas, and the Ural Mountains, which suggests that these regions may be important for modulating weather patterns in East Asia. The climatological mean distribution of GPH at 300 hPa shows troughs located in Eastern Europe and East Asia (Fig. 2b). During the L cases, the Siberian high pressure and Aleutian low pressure were strengthened in the lower-level atmosphere, and the upper-level trough at 300 hPa also intensified (Fig. 2c). This is similar to the pattern of cold waves over the Korean Peninsula in winter, which is consistent with previous studies (Takaya and Nakamura, 2005; Overland et al., 2015; Heo et al., 2018; Sung and Kim, 2020; Overland et al., 2021). This indicates that strong dispersion due to intensified northerly winds can lower PM<sub>10</sub> concentrations. In contrast to the cold wave pattern shown in Fig. 2c, H cases are correlated with a weakening of both the Siberian high pressure and Aleutian low pressure at the lower-level, and with an anticyclonic anomaly in the upper-level of the Korean Peninsula (Fig. 2d). In this pattern, the PM<sub>10</sub> concentration increases as a result of relatively weakened northerly winds and stagnant high pressure in the upper-level atmosphere, as observed in previous studies (Lee et al., 2011; Lee et al., 2018; Kim et al., 2019; Jung et al., 2019).

The circulation patterns during EH cases were noticeably different to that during L and H cases. First, the magnitude of the pressure anomaly

was approximately three times larger than that of the L and H cases. Second, the spatial distribution of the anomalies differed, especially in the North Atlantic and Arctic Oceans, with higher GPH anomalies. This suggests that L and H cases in Korea are dominantly affected by the EAWM scale of circulation, whereas EH cases are dominantly influenced by the Eurasian Continental (EC) scale of circulation. Anticyclonic anomalies (anomalous high pressure) cause the atmosphere to stagnate, resulting in increased pollutant concentrations. The PBLH was much shallower during the EH cases than during the L and H cases (Fig. s1), indicating that local meteorological conditions also provide unfavorable conditions for vertical dispersion. Anticyclonic anomalies are dynamically associated with the descending motion that lowers the PBLH (Zhong et al., 2019). Thus, Fig. 2e is consistent with Fig. s1d. This anticyclonic anomaly and the PBLH anomaly are consistent with the temperature inversion (Fig. s2).

As the atmospheric circulation patterns influencing EH cases have not been thoroughly investigated despite their significant impact on society, we focused on EH cases by analyzing the development of EC-scale circulation patterns, particularly in the Atlantic Ocean and Eurasia. Fig. 3 shows the PM<sub>10</sub> concentration time series five days before and after the occurrence of EH cases. Herein, we defined the first day of 20 EH PM<sub>10</sub> cases ( $\geq 100 \mu\text{g}/\text{m}^3$ ) as the start day (lead-lag = 0). The minimum observed PM<sub>10</sub> concentration during this time series was approximately 44.7  $\mu\text{g}/\text{m}^3$ , which was recorded four days before the extreme event. This value then rapidly increased to approximately 117.4  $\mu\text{g}/\text{m}^3$  (2.6 times) over the next four days, which suggests that PM<sub>10</sub> concentrations in Korea can rapidly increase within just a few days.

As previously discussed, EH cases are associated with EC-scale atmospheric circulation patterns, and PM<sub>10</sub> concentrations in EH cases can rapidly increase within four days. Therefore, to investigate the development of an EC-scale atmospheric circulation pattern that favors for the occurrence of EH cases, we analyzed the composite anomaly from five days before the start day to the day of occurrence (hereafter, -5 days; Fig. 4). Interestingly, on -5 days, EH cases showed a pressure pattern associated with a cold wave as well as the associated anomalous low surface temperature and northerly winds on the Korean Peninsula. As the EH start day approached, the Siberian high pressure and Aleutian low pressure gradually weakened in the lower-level atmosphere. On the EH start day, the anticyclonic anomaly from the northern part of China moved southward to the Korean Peninsula in the upper-level atmosphere. Anticyclonic anomalies cause the atmosphere to stagnate, resulting in increased pollutant concentrations. Therefore, it is important to understand the mechanisms that induce stagnant anticyclones during EH cases. In addition, the surface temperature was generally warmer, and the anomalous southerly wind was dominant. Furthermore, anticyclonic anomalies emerged in the upper- and lower-level atmosphere in the Norwegian Sea and the Scandinavian Peninsula from three to four days before and substantially grows at the occurrence of EH cases, which had a barotropic structure. This anticyclonic anomaly was the most noticeable difference in the atmospheric circulation



**Fig. 8.** (a) Vertical distribution of the prescribed vorticity tendency forcing used in the LBM experiment, and (b) model response of sea level pressure (SLP; shading; units: hPa) and geopotential height at 300 hPa (GPH; contour interval of 10 m; units: m; dashed contours are negative). The green box indicates the vorticity tendency forcing region in the eastern North Atlantic and Northern Europe (ENE; 58.5 °N to 72 °N, 4.5 °W to 30 °E). (For interpretation of the references to colour in this figure legend, the reader is referred to the Web version of this article.)

patterns between the H and EH cases. In this study, the region in which the anticyclonic anomaly noticeably appears was defined as the eastern North Atlantic and Northern Europe anticyclonic anomaly region (ENE; 58.5 °N to 72 °N, 4.5 °W to 30 °E).

### 3.2. Temporal development of North Atlantic forcing

To confirm the development of GPH and TMP in the ENE, we displayed the vertical composite anomaly patterns according to the occurrence days of the EH case (Fig. 5). Notably, four days before the occurrence of EH cases, the anticyclonic anomaly intensified in the upper- and lower-level atmosphere in the ENE, and a strong positive anomaly appeared in the upper-level atmosphere (300 hPa). Moreover, the GPH anomaly in the ENE decayed rapidly one day after the EH start day, which is consistent with the rapid decrease in PM<sub>10</sub> concentrations, as shown in Fig. 3.

We investigated the propagation path of the Rossby wave in the upper-troposphere (300 hPa) during EH cases using a WAF analysis (Fig. 6). The results showed a weak divergence of WAF from the ENE five days before the EH cases. However, WAF diverged into two branches three days before the EH cases, one from the ENE region to Europe and the other to East Asia along the Siberia. Particularly, the second branch reaches the Korean Peninsula. The rapid growth of the anticyclonic anomalies over the ENE likely plays as a source of the Rossby wave that induces the anomalous atmospheric circulation patterns favorable for the EH cases in the Korean Peninsula.

An anticyclonic anomaly in the ENE, which develops strongly in barotropic structures in the upper- and lower-level atmosphere, was associated with the occurrence of EH cases in Korea. Therefore, we confirmed the QG tendency budget for the presence of remote forcing, which affected the strongly developed anticyclonic anomaly in the ENE (Fig. 7). The solution of the QG tendency budget represents the sum of the three forcings (e.g., vorticity advection, temperature advection, and diabatic heating), and the solution is similar to the QG tendency budget calculated using the reanalyzed data. In the upper-level atmosphere of the ENE, the QG tendency showed positive anomalies before the start day of EH cases, but the QG tendency rapidly decreased from the start day of EH cases. In particular, among the three forcings, the QG tendency due to vorticity advection was the most similar to the QG tendency of all forcings, and the QG tendencies due to diabatic heating and temperature advection were found to cancel each other out. Over the ENE, the GPH anomaly developed four days before the EH case (hereafter, -4 days), and the maximum value appeared on days 0 and +1 (Fig. 5). According to the QG tendency budget analysis, the tendency term was positive from -4 to -1 days (0 days for reanalysis data) and gradually decreased from -1 days. While the tendency term maintained a positive value, the GPH anomaly accumulated, which increased the positive GPH anomaly. Therefore, Fig. 7 is generally consistent with Fig. 5. The growth of the forcing that triggered the stationary Rossby wave in the ENE also coincided with the growth of the anticyclonic anomaly over the Korean Peninsula, leading to a maximum at 0 days. These results indicate that the anticyclonic anomaly strongly developed in the upper-level atmosphere of the ENE is dominantly influenced by vorticity advection forcing.

### 3.3. Linear baroclinic modeling experiment

We conducted an LBM experiment to confirm whether the anticyclonic anomaly developed over the ENE induces the EC-scale atmospheric circulation pattern related to EH cases in Korea (Fig. 8). In this experiment, we prescribed the mean tendency of vorticity advection on -4 to -1 days to reproduce the averaged response of the atmospheric circulation over Eurasia. The results of the LBM experiment showed a strongly developed anticyclonic anomaly in the upper- and lower-level atmosphere over the ENE, as well as the development of negative and positive anomalies in Siberia and East Asia, respectively (Fig. 8b). In

addition, the southeasterly winds dominate in the lower-level troposphere over the Korean Peninsula. This lower-tropospheric circulation pattern provides favorable condition for EH cases by weakening the seasonal northwesterly prevailing during the winter (Lee et al., 2020a). This pattern is consistent with the observed pattern shown in Fig. 2e. This suggests that the negative vorticity advection in the ENE induced the observed EC-scale atmospheric circulation pattern, which is favorable for EH cases on the Korean Peninsula.

Vorticity advection in the ENE may have various causes, such as the North Atlantic Oscillation (NAO) and sea surface temperature variability in the North Atlantic (Hurrell and Deser, 2010; Comas-Bru and McDermott, 2014). In particular, NAO is a dominant mode of climate variability in Europe and North America. Therefore, we investigated the potential links between EH cases and the NAO (Fig. s3) and found that the circulation anomalies for the negative and positive NAO phases were not consistent with those of the EH cases. Interestingly, the circulation anomaly of the neutral NAO phase was similar to that of the EH cases (Fig. s3b), indicating that the teleconnection via ENE forcing is effective when the NAO is neutral.

## 4. Conclusion

In this study, we classified the daily PM<sub>10</sub> concentrations into three classes using the median and 95th percentile values: L cases (<50 µg/m<sup>3</sup>), H cases (50–100 µg/m<sup>3</sup>), and EH cases (≥100 µg/m<sup>3</sup>). Our results clearly showed that a strengthening (weakening) EAWM on the Asian monsoon scale influenced the PM<sub>10</sub> concentrations during L (H) cases. Interestingly, EH cases showed strong anomaly patterns on EC-scale atmospheric circulation, which was distinct from that of the L and H cases.

The ENE was the key region affecting the occurrence of EH cases, which corresponds to the North Atlantic Ocean, including the Norwegian Sea and the Scandinavian Peninsula. The rapid growth of the barotropic warming anomaly over the ENE weakens the trough in the upper troposphere over the Korean Peninsula through stationary Rossby waves in the upper-level atmosphere; this induces a shallower PBLH and weakens the wind speed. As a result, the PM<sub>10</sub> concentration during EH cases sharply increased by ~2.6 times from 44.7 µg/m<sup>3</sup> to 117.4 µg/m<sup>3</sup> in just four days. Our findings highlight the importance of understanding the teleconnection between ENE forcing and EC-scale atmospheric conditions for predicting the occurrence of EH cases in Korea and reducing the human health risks of atmospheric pollution.

### Author statement

Jeong-Hun Kim: Conceptualization, Formal Analysis, Investigation, Methodology, Visualization, Writing—Original Draft. Seong-Joong Kim: Conceptualization, Investigation, Methodology, Writing—Review & Editing. Daeok Yun: Investigation, Methodology, Writing—Review & Editing. Maeng-Ki Kim: Conceptualization, Investigation, Methodology, Writing—Review & Editing. Joo-Hong Kim: Methodology, Writing—Review & Editing. Joowan Kim: Methodology, Writing—Review & Editing. El Noh: Methodology, Writing—Review & Editing.

### Declaration of competing interest

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

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## Appendix A. Supplementary data

Supplementary data to this article can be found online at <https://doi.org/10.1016/j.envpol.2021.118051>.

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